

Benthic jellyfish dominate water mixing in mangrove ecosystems

David M. Durieux^{a,1}, Kevin T. Du Clos^{a,2}, David B. Lewis^a, and Brad J. Gemmell^a

^aDepartment of Integrative Biology, University of South Florida, Tampa, FL 33620

Edited by Mark Denny, Stanford University, Pacific Grove, CA, and accepted by Editorial Board Member Simon A. Levin June 16, 2021 (received for review January 4, 2021)

Water mixing is a critical mechanism in marine habitats that governs many important processes, including nutrient transport. Physical mechanisms, such as winds or tides, are primarily responsible for mixing effects in shallow coastal systems, but the sheltered habitats adjacent to mangroves experience very low turbulence and vertical mixing. The significance of biogenic mixing in pelagic habitats has been investigated but remains unclear. In this study, we show that the upside-down jellyfish Cassiopea sp. plays a significant role with respect to biogenic contributions to water column mixing within its shallow natural habitat (< 2 m deep). The mixing contribution was determined by high-resolution flow velocimetry methods in both the laboratory and the natural environment. We demonstrate that Cassiopea sp. continuously pump water from the benthos upward in a vertical jet with flow velocities on the scale of centimeters per second. The volumetric flow rate was calculated to be 212 L·h⁻¹ for average-sized animals (8.6 cm bell diameter), which translates to turnover of the entire water column every 15 min for a median population density (29 animals per m²). In addition, we found Cassiopea sp. are capable of releasing porewater into the water column at an average rate of 2.64 mL·h⁻¹ per individual. The release of nutrient-rich benthic porewater combined with strong contributions to water column mixing suggests a role for Cassiopea sp. as an ecosystem engineer in mangrove habitats.

water mixing | jellyfish | *Cassiopea* | mangroves | particle image velocimetry

oastal mangrove habitats exhibit exceptionally high productivity and provide many ecosystem services, including control of flooding, sedimentation, and nutrient input in surrounding areas (1). In addition, these habitats act as critical nursery habitat for early life stages of a wide range of commercially and ecologically important species of fish and crustaceans (2, 3) and as adult habitat for many others (3). Mixing, which stirs gradients and homogenizes the water column, is a critical process which regulates important processes such as productivity (4) and benthic-pelagic coupling (5), prey encounter rate (6), and gas exchange (7). Thus, it is important to identify and understand the major sources of mixing within these coastal environments. Mixing forces are generally physical in origin, such as wind or tides, but, in some cases, they can originate from biotic sources. Biogenic mixing may be of particular importance in mangrove ecosystems, as these habitats experience comparatively little physical mixing due wind or tidal flow (8, 9).

Previous studies have explored the role of biogenic mixing contributions of several taxa, including pelagic jellyfish, fish, and krill in the water column (10, 11), as well as benthic bivalves (12). These studies suggest the potential for nontrivial levels of animal-mediated mixing even in unsheltered environments, although others suggest a much lower level of mixing contribution (13) in the global oceans relative to processes such as internal waves. Little work is available regarding such effects in shallow (< 2 m) habitats, however. One process by which internal waves might be produced is through interactions between tidal

PNAS 2021 Vol. 118 No. 30 e2025715118

flow and topography in stratified waters, and these waves have been shown to interact with coastal ecosystems (14). However, in the much shallower, quiescent conditions of mangrove lagoons, these conditions do not appear to be present.

While the turbulent mixing in shallow coastal ecosystems is often driven by wind and wave action (15), mangrove ecosystems provide shelter from wind and waves (9) and dampen mixing from high-velocity tidal currents, which tend to run, primarily, in deeper channels within mangrove ecosystems (16). Therefore, physical mixing is naturally low in *Cassiopea* sp. habitat, implying that animal-mediated mixing could have more substantial impacts in these environments.

The upside-down jellyfish, *Cassiopea* sp., has a nearly circumtropical distribution (17). They are found in shallow, low-energy coastal environments, often dominated by the red mangrove, *Rhizophora mangle* (18), and have been previously documented at densities up to 31 animals per m² or 20% benthic coverage (19). In our field site, they were encountered at water depths of < 2 m. Their natural range is extensive, and is further spreading due to their anthropogenic introduction as an invasive species (20–22) and through natural range expansion due to rising ocean temperatures (21). In addition, the size and abundance of *Cassiopea* sp. have increased near anthropogenically disturbed habitats, where nutrients tend to be elevated (23). These factors

Significance

Water mixing is critical for aquatic life and mangrove habitats are vital nurseries for many species. However, these sheltered habitats experience little water mixing. *Cassiopea* sp. jellyfish continuously produce a vertical current capable of transporting hundreds of liters of seawater per hour. They live in shallow subtidal regions and the flows they generate have the potential to result in turnover of the entire water column up to every 15 min for an average population. Additionally, *Cassiopea* sp. can expedite the transport of nutrient-rich water from sediments into the water column. Thus, *Cassiopea* sp. represent a previously unrecognized ecosystem engineer within these shallow coastal environments that can affect primary productivity and nutrient distribution and alter new habitats as their range is expanding.

The authors declare no competing interest.

This article is a PNAS Direct Submission. M.D. is a guest editor invited by the Editorial Board.

Published under the PNAS license.

This article contains supporting information online at https://www.pnas.org/lookup/suppl/ doi:10.1073/pnas.2025715118/-/DCSupplemental.

Published July 23, 2021.

Author contributions: D.M.D., K.T.D.C., and B.J.G. designed research; D.M.D., K.T.D.C., and B.J.G. performed research; K.T.D.C., D.B.L., and B.J.G. contributed new reagents/analytic tools; D.M.D., K.T.D.C., and B.J.G. analyzed data; D.M.D. wrote the paper; and D.M.D., K.T.D.C., and B.J.G. edited versions of the paper.

¹To whom correspondence may be addressed. Email: daviddurieux@mail.usf.edu.

² Present Address: Institute of Ecology and Evolution, University of Oregon, Eugene, OR 97403.



Fig. 1. (*A*) Laboratory and (*B*) in situ PIV of a representative Cassiopea sp. jet. White arrows represent the velocity vectors, while colored areas represent the flow with either an upward (red) or downward (blue) component. (C) A cluster of animals produces multiple adjacent jets.

make *Cassiopea* sp. a potential indicator species for a variety of anthropogenic disturbances.

Unlike most other jellyfish, *Cassiopea* sp. exhibit an epibenthic lifestyle, with their bell on the substrate and oral arms facing upward. They pulse their bells nearly continuously, producing a vertical jet (Fig. 1) (24–26); however, the extent of this vertical jet and volumetric flux have not been fully quantified. The direction of flow and volumetric fluxes may be particularly noteworthy given the fact that, in the absence of *Cassiopea* sp., there is a downward flux of nutrients, with the sediments acting as a nutrient sink in sheltered mangrove habitats (27). Thus, an upward pump at the benthos may serve to alter this flux and also interact with interstitial porewater, pulling a fraction of this water into the water column (26) and potentially affecting local productivity, as has been seen where benthic nutrients are released by bioturbation (28) or wind events (4).

In this study, we quantify the large-scale flow features of individual *Cassiopea* sp. as well as flows created by small groups both in the laboratory and in the field in order to test the hypothesis that range expansion and alterations in population densities of these benthic jellyfish could significantly increase mixing in otherwise quiescent habitat through the vertical excurrent jet of their feeding current. We discuss the results in the context of environmental mixing in mangrove ecosystems and examine the role of biological pumping at the surface of the benthos with respect to the liberation of interstitial porewater from the sediments.

Results

Field Population. At our study site in the Florida Keys, the density of the studied population of *Cassiopea* sp. ranged from 3 animals per m² to 97 animals per m², with a median density of 29 animals per m² (Fig. 2*C*) (29, 30). The sizes of these animals followed a Poisson distribution (*SI Appendix*, Fig. S1, χ^2 Test for Independence, *P* < 0.001) with a mean of 11.3-cm oral arm diameter



Fig. 2. (*A*) Vertical flux of water increases exponentially and follows a power law relationship with animal size ($R^2 = 0.732$, n = 9). (*B*) The relative vertical flux, such that a value of one is equal to the amount an individual pumps in the absence of neighbors, decreases as the number of neighbors increases ($R^2 = 0.74$, n = 12). Points represent individual jellyfish, while curves show the best fit model. (*C*) Predicted time to turn over the column of water above a square-meter patch of *Cassiopea* sp. assuming a water depth of 1 m, which was typical in our field site, relative to the population density in that patch (line). Frequency of different population densities (bars) observed is shown as a histogram of $1-m^2$ plots.

(OAD) and 8.6-cm resting bell diameter (RBD), and there is a linear relationship between OAD and RBD (*SI Appendix*, Fig. S2 and Eq. S1; n = 50, $R^2 = 0.9$). We calculated that animals with the observed average RBD of 8.6 cm have a wet weight (WW) of 56.5 g (*SI Appendix*, Fig. S2 and Eq. S2; n = 50, $R^2 = 0.99$) and a dry weight (DW) of 3.2 g (*SI Appendix*, Fig. S2 and Eq. S3; n = 31, $R^2 = 0.98$).

Quantitative Description of the Vertical Jet. Particle image velocimetry (PIV) was performed on nine Cassiopea sp. ranging in size from 4.8-cm RBD to 14.5-cm RBD (31). Analysis of the shape of the jet produced by these animals showed that the jet diameter (D_i) , relative to RBD, spreads linearly as height (Z) increases (SI Appendix, Fig. S2 and Eq. S4; linear regression: $R^2 = 0.64$, p < 0.001). Maximum velocity at each height correlated by a power function to height normalized to bell diameter (SI Appendix, Fig. S2 and Eq. S5; $R^2 = 0.33$). Our regression predicts a mean velocity of 20 mm \cdot s⁻¹ at benthos level regardless of animal size, but the rate of water speed slowing above this height decreased as animal size increased. At a height of one bell diameter above the animal, Re =1,091 based on jet diameter. An extrapolation assuming calm conditions in the field shows that, for the average-sized animal (RBD = 8.6 cm), the jet would slow to 3.86 mm \cdot s⁻¹ at a height of 2 m. Since Cassiopea sp. in this study were found only between 0.5 and 2 m deep, it is assumed that the jets will reach the surface in their natural habitat. In situ PIV confirmed that wild Cassiopea sp. in the field produced maximum vertical flow velocities similar to our laboratory-based velocity measurements.

Flux and Turnover Time. At a relative distance of two bell diameters above the animals, vertical velocity was found to be independent of bell diameter (Spearman's rank correlation, $\rho = 0.233$, p = 0.551), as was expected from the observation of jet shape. Because jet diameter was found to correlate directly with bell diameter (Spearman's rank correlation, $\rho = 0.9$, p = 0.002), and velocity was independent of bell diameter, vertical flux of water (Q) was found to follow a power law function relative to RBD, with a power of 2.23 (Fig. 2A and SI Appendix, Fig. S2 and Eq. S6; n = 9, $R^2 = 0.74$). The relationship can also be expressed as weight-specific flux (F) in terms of DW (SI Appendix, Fig. S2 and Eq. S7). Thus, an average-sized Cassiopea sp. moves 212 L·h⁻¹ of seawater, a mass-specific volumetric flow rate of 62.9 L·h⁻¹·g⁻¹ DW.

Since *Cassiopea* sp. form aggregations in nature, a correction factor was needed to account for flow-related interactions between individuals. The jet of an individual *Cassiopea* sp. was

examined under conditions with different numbers of neighbors (32), and the flux was found to decrease with the addition of each neighbor (Fig. 2B and SI Appendix, Fig. S2 and Eq. S8), where neighbors were defined as being in physical contact with each other. Based on photographic transects of the field population in the Florida Keys, at the median density of 29 animals per m^2 , each animal will have an average of 1.23 neighbors (SI Appendix, Fig. S2 and Eq. S9). Assuming a water depth of 1 m and a vertical jet that reaches the surface, the turnover time for the water column was calculated. A single animal would turn over the water column over a square meter of benthos in 4.7 h. At the observed median population of 29 animals per m², the resulting flux is 3,980 L \cdot h⁻¹, and the calculated turnover time is reduced to 15 min. Turnover time decreases as population density increases, to a minimum of 11.8 min at a density of 64 animals per m^2 (Fig. 2C), at which point turnover time begins to increase due to increased interference between neighbors.

In Situ Flow Measurement. In the Lido Key site, in situ PIV demonstrated that *Cassiopea* sp. increase vertical transport of water in these habitats. Mean vertical velocity increased from $0.12 \text{ mm} \cdot \text{s}^{-1}$ downward ($\pm 0.39 \text{ mm} \cdot \text{s}^{-1}$ SD, n = 7) in the absence of jellyfish to $4.01 \text{ mm} \cdot \text{s}^{-1}$ upward ($\pm 1.52 \text{ mm} \cdot \text{s}^{-1}$ SD, n = 5) (Fig. 3). This difference was determined by an unpaired *t* test to be highly significant (t = 5.8487, df = 10, P = 0.0002). The horizontal component of the flow showed a smaller, but still significant change, reducing from 19.82 mm \cdot \text{s}^{-1} in the absence of jellyfish to 12.97 mm \cdot \text{s}^{-1} when a jellyfish was present (t = 3.2021, df = 10, P = 0.0095). Turbulent kinetic energy (TKE) dissipation rates were $3.5 \times 10^{-6} \text{ m}^2 \cdot \text{s}^{-3}$ ($\pm 1.8 \times 10^{-6} \text{ m}^2 \cdot \text{s}^{-3}$ SD, n = 4). In the presence of jellyfish, this increased to $6.5 \times 10^{-6} \text{ m}^2 \cdot \text{s}^{-3}$ ($\pm 4.3 \times 10^{-6} \text{ m}^2 \cdot \text{s}^{-3}$ SD, n = 8). This increase, while measureable, was not statistically significant (two-sample *t* test, P = 0.2).

To confirm that these habitats are quiescent and protected from wind-driven wave action, a comparison between wind waves between the mangrove habitat and the adjacent Sarasota Bay showed a reduction in wave height from $5.4 \pm$ 2.4 cm (N = 10) in the bay to 0.07 ± 0.07 cm (N = 10) in the mangroves (*SI Appendix*). In addition, the horizontal velocity measured during our PIV experiment was 19.82 cm·s⁻¹ in the absence of jellyfish. Since this experiment was performed near peak tidal flow in order to avoid resampling water masses, this value must be near the upper limit for tidal velocities in this habitat. In the absence of *Cassiopea* sp., the water column was weakly temperature stratified (*SI Appendix*, Fig. S4).



Fig. 3. Representative time-averaged vertical flow component in the (*A*) presence and (*B*) absence of *Cassiopea* sp. at Lido Key, FL, as measured using particle image velocimetry over 1 s. Blue indicates downward flow; red indicates upward flow. Red line above PIV image represents the spatially averaged SD in vertical velocity over the period of the time-averaged image. (*C*) Quantitative analysis shows that the presence of *Cassiopea* sp. significantly increased vertical flow from $-0.12 \text{ mm} \cdot \text{s}^{-1}$ to 4.01 mm $\cdot \text{s}^{-1}$. (Unpaired *t* test, df = 10, p = 0.0002).

Interstitial Water Release. Interstitial water release was quantified by measuring changes in concentrations of fluorescein in water over fluorescein-laced sand. Fluorescein dye concentrations increased visibly in the trials that included a jellyfish, but

https://doi.org/10.1073/pnas.2025715118

not in controls (33). The ratio of absorbance at 494 nm between t = 3 h and t = 2 h increased from 1.02 (± 0.07 SD, n = 5) to 1.15 (± 0.16 SD, n = 9) in the presence of a jellyfish. From this absorbance, we calculated the average flow attributed to a

4 of 8 ∣

PNAS

single jellyfish, normalized to an RBD of 8.6 cm, which amounted to 2.64 mL·h⁻¹ (± 2.1 mL·h⁻¹ SD, n = 9) of interstitial water released into the water column (Fig. 4). Trials without jellyfish had no apparent fluorescein release. The flow rate in the presence of jellyfish was found to vary significantly from zero (one-sample t test, P = 0.0055). Interstitial water release rate increased linearly with bell diameter (Fig. 4, n = 9, $R^2 = 0.3622$).

Discussion

In many coastal ecosystems, wind and tidal flow are important mixing forces (15), but the ecosystems inhabited by Cassiopea sp. are sheltered by dense stands of *Rhizophora* mangrove trees, reducing wind-water interaction in these lagoons and estuaries (9). In addition, *Cassiopea* sp. dwell on flats at depths of < 2m, outside of the tidal channels where most tidal flushing in these habitats occurs (8). Together, these two habitat features result in a system with comparatively little physical mixing. It is possible that *Cassiopea* sp. utilize these habitats due to a risk of being swept away by stronger flow and thus unable to settle on the bottom. Evidence of the quiescent nature of this habitat is given by the presence of a fine-grain benthic substrate in the areas where *Cassiopea* sp. are abundant (34), and our field measurements found TKE dissipation rates comparable to the low end of previously reported measurements (35) from mangrove habitat. This confirms that environmental flows in which



Fig. 4. (A) Vertical flux of interstitial water out of the sediment attributed to a single Cassiopea sp., normalized to an RBD of 8.6 cm (the average observed in the field), compared to the rate due to diffusion in the experimental container. Interstitial water release due to diffusion was negligible, while the average Cassiopea sp. released 2.6 mL·h⁻¹. (B) Vertical flux of interstitial water relative to bell diameter. As animal size increased, the interstitial flow increased as well (n = 9, $R^2 = 0.3622$).

Cassiopea sp. are found are low, with little turbulent mixing. As such, the cumulative effect of a population of Cassiopea sp. has the potential to provide a large contribution to mixing relative to nonbiogenic factors, which would have important ecological implications.

Our measurements and calculations show that the average animal has the potential to produce a jet of continuous flow on the order of several meters tall, which would certainly affect the entire depth of the water column (0.5 m to 2 m) in the shallow ecosystem where Cassiopea sp. occur. This flow occurs at transitional Reynold's numbers, and this is reflected in the slowing rate of the vertical jet. Our results indicate that an individual average-sized animal can move 212 L·h⁻¹ of seawater upward in its vertical jet (Fig. 2A) at velocities of a similar order of magnitude to those found in prior work (25). It is known, however, that multiple adjacent jets can interfere with one another, either by redirecting momentum radially and reducing axial momentum (36) or, as observed in the asynchronous pulsing of jets in colonial salps, by producing a more steady flow at the possible expense of thrust (37). After factoring in the effect of neighboring animals, our model predicts that the median population density observed in our field site (29 animals per m^2) can turn over a 1-m deep column every 15 min (Fig. 2C). This biogenic contribution to environmental mixing is higher than any other known epibenthic species (Table 1).

For comparison (Table 1), the mussel Mytilus edulis has been observed to pump 6.4 $L \cdot h^{-1}$ to 11.9 $L \cdot h^{-1}$, and the oyster Crassostrea virginica has been calculated to pump 2.61 $L\cdot h^{-1}$, in both cases, two orders of magnitude below the rates measured in Cassiopea sp. in this study. To account for differences in body mass and body water content, DW-specific filtration rates can also be compared (Table 1). The mussel M. edulis pumps 5.94 6.4 $\text{L}\cdot\text{h}^{-1}\cdot\text{g}^{-1}$ DW, and the eastern oyster *C. virginica* moves 6.4 $\text{L}\cdot\text{h}^{-1}\cdot\text{g}^{-1}$ DW. By comparison, *Cassiopea* sp. exhibit volumetric flow rates an order of magnitude higher than these bivalves by DW. Therefore, per unit of carbon, Cassiopea sp. act as much more effective water pumps than do bivalves. In terms of WW, Cassiopea sp. have weight-specific filtration rates that are also an order of magnitude above those of bivalves (Table 1). The weight-specific pumping rate of the benthic tunicate Cionia intestinalis is at a level closer to that of Cassiopea sp. However, due to the much smaller size of C. intestinalis, they move a lower magnitude of water than *Cassiopea* sp. overall.

Of course, animals such as oysters and mussels can occur at much higher densities than Cassiopea sp. do, but their relatively small individual flow is not fully compensated for in terms of numbers. High densities of over 700 animals per m² have been reported for both M. edulis (45) and C. virginica (38). Assuming adult [0.3 g DW (38)] oysters and additive contributions (a likely overestimate), this would produce approximately 1,176 $L \cdot h^{-1} \cdot m^{-2}$ flow. For mussels (42) and the same assumptions, we calculate a flow of 3,443 $L \cdot h^{-1} \cdot m^{-2}$. The tunicate *C. intestinalis* was found to pump water at rates of 0.1 $L h^{-1}$ to 1 $L h^{-1}$ (43). This species, has been recorded at densities of over 1,000 individuals per m^2 (46) and has a population filtration rate on the order of 1,000 L h⁻¹ m⁻². Our median population of *Cassiopea* sp. (29 animals per m²) moved 3,980 L·h⁻¹·m⁻², while peak densities observed in this study moved 4,569 $L \cdot h^{-1} \cdot m^{-2}$. This translates to peak bulk flow estimates from Cassiopea sp. that are 36% higher than peak values for mussels, 386% higher than oysters, and 457% higher than benthic tunicates. Thus, Cassiopea sp. populations, despite their lower population densities, are responsible for greater amounts of water transport than are bivalve reefs or tunicate colonies and are occurring in an area with comparatively

COLOG

Durieux et al.

Table 1. Estimated filtration rate per individual, as well as WW- and DW-specific filtration rates for *Cassiopea* sp. (assuming a density of 29 animals per m²), *C. virginica*, *M. edulis*, and the benthic tunicate *C. intestinalis*

	Filtration rate (L·h ⁻¹)	WW-specific filtration rate $(L \cdot h^{-1} \cdot g^{-1})$	DW-specific filtration rate $(L \cdot h^{-1} \cdot g^{-1})$	Population bulk transpor (L·h ⁻¹ ·m ⁻²)
Cassiopea sp. (Current study)	137	2.43	42.89	4,569
C. virginica	2.61 (38, 39)	0.71 (39, 40)	6.41 (39)	1,176
M. edulis	11.9 (12)	0.80 (12, 41, 42)	5.94 (12, 42)	3,443
C. intestinalis	0.1 to 1 (43)	1.16 to 1.57 (43, 44)	24 (44)	1,000

Rates for *Cassiopea* sp. are consistently an order of magnitude larger than those of bivalves. Citations refer to sources of filtration rates and size conversion factors used to calculate these values.

little physical mixing, in comparison to *M. edulis* which inhabits much more physically dynamic conditions (47). In addition, *Cassiopea* sp., unlike other biological pumps that are smaller, are capable of producing jets which reach the surface of the water (*SI Appendix*, Fig. S2 and Eq. S5). Other, smaller pumps are much more limited in the extent of their influence, with *M. edulis* jets decomposing within 5 cm of the bottom (48). The result of this difference in extent is that *Cassiopea* sp. is able to turn over the entire water column, while bivalves and other smaller organisms mix only the lower strata.

The fact that median *Cassiopea* sp. populations are capable of turning over the entire shallow water column in 15 min has implications for several ecosystem processes. For one, this vertical mixing could disrupt the temperature profile observed in our site (SI Appendix, Fig. S4), homogenizing the water column. An increase in mixing, especially in upward vertical transport, could lead to a decrease in sedimentation in low-energy habitats. Particulates in mangrove habitats flocculate, increasing particle size, but the low density of these flocs reduces sedimentation rate (49). Cassiopea sp. themselves may contribute to flocculation by the production of mucus (19), which traps suspended particles (49). On the other hand, these particles may be further prevented from settling by addition of mixing energy from Cassiopea sp., thereby increasing the export of nutrients from mangrove ecosystems to neighboring systems such as coral reefs and seagrass beds. Further investigations would be required to determine the effects of Cassiopea sp. on sediment transport. Another important consideration for Cassiopea sp.-mediated biogenic mixing is that the strongest flows are generated near the benthos. This is in contrast to wind-driven mixing in which the strongest flows are at the surface, and may have important implications for productivity in these habitats. For example, a mesocosm study where mixing was initiated close to the bottom, as is the case with Cassiopea sp., found that faster turnover resulted in longer-lasting algae blooms and higher chlorophyll concentrations than at lower mixing rates (6), likely because of the increased availability of nutrients from the sediment to phytoplankton cells. In addition, the accelerated exchange of water near the benthos to the surface in the presence of Cassiopea sp. may affect oxygen and CO2 exchange rates at the surface, affecting carbonate pH buffering, respiration rates, and primary productivity throughout the water column.

It has been previously demonstrated that *Cassiopea* sp. can transport interstitial porewater in sediments from several centimeters deep in the benthos to the water column (26), and, while the maximum velocity of this flow was similar to those in this study at 2 cm·h⁻¹, volume fluxes were not measured, and

thus the potential impact on the local ecosystem is unknown. In Gulf of Mexico sediments, increasing the advective benthic flow produced an increase in denitrification rates (50), so it is possible that, in areas where Cassiopea sp. are abundant, they may also contribute to increased denitrification. Additionally, increased flow alters the redox gradient and microbial diversity in benthic sediments (51). Our results show that while, over the course of the experiment (3 h), diffusion released negligible amounts of benthic porewater, an average-sized animal (8.6 cm) would release approximately 2.6 mL h^{-1} of porewater (Fig. 4), and the median population of 29 animals per m² would therefore release 1,800 mL·d⁻¹·m⁻². Given that these animals tend to move slowly over the bottom (26), thereby encountering new sediments regularly, and given the reported phosphate and nitrogen concentrations in the sediment (52) at Long Key, FL, the action of *Cassiopea* sp. could release roughly $0.34 \text{ mg} \cdot d^{-1} \cdot m^{-2}$ of NO₃, 2.35 mg $\cdot d^{-1} \cdot m^{-2}$ of NH₄, and 1.03 mg $\cdot d^{-1} \cdot m^{-2}$ of PO₄ at median population densities. Since these released nutrients will be mixed into the water column upon becoming entrained in the vertical jets and assuming a 1-m water depth, this represents an increase of 29% in water column NH₄, a 6% increase in PO₄, and a 0.5% increase in NO3 relative to reported water column nutrient concentrations from these habitats (52). The addition of the mixing contribution produced by Cassiopea sp. into a habitat with low environmental water mixing, as well as a substantial release of interstitial water leading to an additional source of nutrients in this ecosystem, make it likely that Cassiopea sp. are previously unrecognized ecosystem engineers within mangrove habitats.

We have demonstrated that *Cassiopea* sp. play a key role in biogenic mixing in quiescent mangrove habitats, where other significant physical mixing processes are absent. This, along with the potential effects of *Cassiopea* sp. on nutrient flux, lead to several future lines of study. In particular, it wold be interesting to determine the in situ effects of this flow on phytoplankton productivity and on turbidity. In addition, demonstrating any effects of interstitial water flow on the benthic environment, in terms of oxygen profiles and benthic communities, would further support the claim that this species is an ecosystem engineer.

Materials and Methods

Field Population. The field study site on Long Key, FL (24.527 N, -80.814 W), was sheltered by red mangroves (*R. mangle*) and held a population of *Cassiopea* sp. at depths ranging from 0.5 m to 2 m deep. The benthic habitat where *Cassiopea* sp. were most abundant consisted of very fine sediment, surrounded by seagrass (*Thalassia* and *Syringopodium*).

The population structure of a natural *Cassiopea* sp. population was quantified during December 2016, by setting two transects through the patch, 10 m apart, parallel to one another in a roughly east-west orientation. Quadrats of 1 m² were sampled every 1.5 m on each transect to reduce investigator bias in selecting animals for measurement. Each quadrat was photographed from above, and the OAD of each animal was measured from these images by measuring the longest observable distance from one oral arm tip to another oral arm tip. During the summer of 2017, 50 animals were collected and imaged from above and from the side. The RBD (used for lateral imaging) was measured during the resting phase between bell contractions using Image image analysis software (53, 54), and was correlated to the OAD, WW, and DW by linear regression to allow conversion between measurements. Based on these parameters, populations of *Cassiopea* sp. at different densities were calculated.

Vertical Jet Imaging. To determine the volumetric flow rate of vertical water movement, the vertical jet must first be described quantitatively. *Cassiopea* sp. of various sizes were collected from the Keys Marine Laboratory on Long Key, FL, during the months of December 2016 and August 2017, and transported in collected seawater back to the University of South Florida, in Tampa, FL, where they were housed in a 300-L closed-loop aquarium system. The animals were kept in artificial seawater mixed to a salinity of 35% with Instant Ocean aquarium salt, over a substrate of aragonite sand and high-intensity metal halide lighting on a 12:12 light cycle.

The imaging setup consisted of a 45 cm \times 45 cm \times 45 cm aquarium, filled with artificial seawater. The water was seeded using 10-µm reflective hollow glass spheres for PIV. An Edgertronic high-speed camera filming at 50 frames per s provided a field of view ca. 30 cm \times 30 cm. Two 2-W continuous wave diode-pumped solid state (DPSS) lasers (wavelength = 520 nm), spread through cylindrical lenses to produce narrow light sheets, were staggered one above the other to illuminate a single coronal plane across the entire field of view. One jellyfish at a time (n = 9) was placed on the bottom in the center of the aquarium, such that the laser sheet crossed the center of the animal. After allowing it to settle for about 10 min, 30 s of video were recorded at 50 frames per s. At least five such image sequences were made, and the three with the most similar pulse rates were retained for analysis using the LaVision software package, producing a PIV time average over the 30 s. PIV was processed with interrogation windows between 48 and 64 pixels, at 50% overlap.

A custom MATLAB script was used to identify and measure the vertical jet from this time average. The location of the jet was defined as the region where the vertical velocity was greater than 0.5 mm·s⁻¹. This region was used to calculate the jet diameter. Due to the asymmetric meandering of the jet, volumetric flux was calculated by using a variant of the washer method. We integrated vertical velocity between the edges of the jet, using half-washers on each side of the position of maximum upward velocity. Because of the high degree of irregularity in *Cassiopea* sp. jets, the three image sequences of each animal were combined to create a single representative jet by taking the median values for jet diameter, jet area, maximum and average velocity (Vz), and vertical volumetric flux (Q) at each height, for use in further analysis.

This method produces an artificial tapering of the jet shape toward its highest end due to the fixed velocity threshold, since the maximum velocity of the jet approaches the minimum threshold as it slows with increasing height. To account for this, data above the height of the maximum measured jet diameter were excluded. In addition, the development region of a turbulent jet does not follow the same patterns as the fully developed region (55). To take this into account, the lowest data points were excluded so that the data series always began at a local maximum in Vz.

For confirmation of the laboratory results, in situ PIV was performed at the field site. A Nikon D750 digital single lens reflex (DSLR) camera at 50 frames per s in an underwater housing was lined up with the bottom near an individual *Cassiopea* sp. with a green laser, spread through a cylindrical lens, also housed in a water-tight case. PIV analysis using the naturally occurring particles was performed using DaVis, and the time-averaged vertical velocities were compared to those measured in the laboratory experiment.

Quantitative Description of the Vertical Jet. Upward Vz was expected to show an inverse relationship (55) with height (*Z*) relative to RBD (*SI Appendix*, Fig. S2 and Eq. S10). In addition, the jet diameter (*D*₁) relative to RBD was hypothesized to increase linearly over this region (55) (*SI Appendix*, Fig. S2 and Eq. S11). Vertical volumetric flux (*Q*) was expected to increase linearly with height, due to entrainment of water into the turbulent jet, and increase linearly with animal size (55). The resulting equation is a function of average velocity and jet diameter (*SI Appendix*, Fig. S2 and Eq. S12). These relationship were tested using the Nonlinear Regression tool in SPSS statistical software version 23.

To confirm that *Cassiopea* sp. in the field produce the same vertical jet as animals in a captive setting, in situ PIV was performed on an animal at the Keys Marine Laboratory, and the parameters of this jet were compared to those captured in the laboratory.

Weight-specific volumetric flow rates were calculated by calculating the expected flux of the animals used for morphological measurements, and then expressing this flux in terms of the WWs and DWs of those animals.

Interactions between Multiple Animals. Because *Cassiopea* sp. are often found in dense aggregations, we determined the degree of interference between adjacent animals, using the same PIV design. A single animal was placed in the filming vessel and then surrounded by an increasing number of neighbors, each in direct contact with the bell of the study animal.

- A. J. Gilbert, R. Janssen, Use of environmental functions to communicate the values of a mangrove ecosystem under different management regimes. *Ecol. Econ.* 25, 323–346 (1998).
- S. Das, Ecological restoration and livelihood: Contribution of planted mangroves as nursery and habitat for artisanal and commercial fishery. World Dev. 94, 492–502 (2017).
- A. I. Robertson, N. C. Duke, Mangroves as nursery sites: Comparisons of the abundance and species composition of fish and crustaceans in mangroves and other nearshore habitats in tropical Australia. *Mar. Biol.* 96, 193–205 (1987).

The vertical velocity and diameter of the original animal's jet were calculated as before, allowing us to determine the degree of inhibition caused by the addition of neighbors. This was compared to the number of neighbors expected in wild populations, determined from the same transect photos used to determine average animal size and population density.

Turnover Time. The time needed for a *Cassiopea* sp. population to turn over the water column above it, where the jet reached the surface, was calculated by taking the volume of water above a $1-m^2$ patch of a hypothetical *Cassiopea* sp. population and calculating the vertical volumetric flux for each animal in that population. The sum of the vertical flux rates (Q) of the animals in this hypothetical patch is divided by the volume of water, giving the turnover rate (*SI Appendix*, Fig. S2 and Eq. S13).

In Situ Flow Measurement. Flow velocities were measured around a population of *Cassiopea* sp. at Lido Key, FL (27.303 N, -82.566 W) using an in situ PIV apparatus recording at 60 frames per s. The PIV plane was aligned with the direction of tidal flow, and the water column above and downstream of a single Cassiopea was recorded over a period of several minutes. The animal was then removed, and recording continued. After PIV vectors were computed, 10-s time averages were taken of the vectors across the entire field of view, spaced apart, temporally, by a time period determined to allow the tidal current to carry the previously measured water mass out of the field of view. This was done to prevent resampling of the same water mass. TKE dissipation rates *ca*. 15 cm above the bottom near patches of *Cassiopea* sp. were calculated from the PIV data using published methods (56), assuming homogeneous and isotropic conditions. Additional confirmatory methods utilized acoustic Doppler velocimetry (57) (*SI Appendix*).

Interstitial Water Release Rates. We labeled play sand by mixing it with a solution of fluorescein in artificial seawater until the sand was uniformly damp. A smooth layer of labeled sand 2.5 cm to 3 cm deep was added to the bottom of an 8-L plastic bucket in a larger water bath, capped with a ca. 1-cm-deep layer of clean play sand mixed with seawater at a rate of 500 mL of seawater to 2 L of dry sand, which delayed leaching of fluorescein into the water column by about 30 min. The bucket was then filled with 4 L of artificial seawater without disturbing the sand layers. Immediately after filling, we took samples of interstitial and column water. A Cassiopea sp. specimen was then placed on the sand in the center of the bucket. On each day of experimentation, an additional trial was performed without a jellyfish to measure diffusion rates in the absence of animals. At 2 and 3 h, the water column was sampled, and the absorbance of this water at 494 nm was compared to a dilution curve of the interstitial water sample. The relative fluorescein concentrations at each time point were corrected by subtracting both the measured concentrations of the starting water sample for the same trial (to control for unintentional fluorescein released during setup) and the concentration of the control bucket at the same time point (to control for diffusion). The change in these corrected values between 2 and 3 h was then used to calculate the rate of interstitial water release into the water column.

Data Availability. Video and data sheets have been deposited in FigShare (https://figshare.com/articles/dataset/PIV_videos_of_clusters_of_Cassiopea_vertical_jet/9754916, https://figshare.com/articles/dataset/PIV_videos_of_ Cassiopea_vertical_jets/9753149, https://figshare.com/articles/dataset/ Interstitial_Porewater_Release/9701534, https://figshare.com/articles/ dataset/Morphological_Measurements_KML_2017/9700406, https://figshare. com/articles/dataset/KML_2016_Population_Transects/9700370, and https:// figshare.com/articles/dataset/Lido_Key_ADV_Data/9696272).

ACKNOWLEDGMENTS. We acknowledge the following people who assisted with specimen collection and experimentation (in alphabetical order): Olivia Blondheim, Alicia Durieux, Christian Fender, Olivia Hawkins, Kathrene Lo, Gabrielle Scrogham, and Nils Tack. Additionally, we thank the Keys Marine Laboratory for use of their facilities. Funding for this project was provided by grants from the NSF (CBET-1511996, IDBR-1455471, and OCE-1829945 to B.J.G.).

- D. Lawrence et al., Wind events and benthic-pelagic coupling in a shallow subtropical bay in Florida. Mar. Ecol. Prog. Ser. 266, 1–13 (2004).
- J. D. Ackerman, M. R. Loewen, P. F. Hamblin, Benthic pelagic coupling over a zebra mussel reef in western Lake Erie. *Limnol. Oceanogr.* 46, 892–904 (2001).
- J. E. Petersen, L. P. Sanford, W. M. Kemp, Coastal plankton responses to turbulent mixing in experimental ecosystems. *Mar. Ecol. Prog. Ser.* **171**, 23–41 (1998).
- C. J. Zappa et al., Environmental turbulent mixing controls on air-water gas exchange in marine and aquatic systems. Geophys. Res. Lett. 34, L10601 (2007).

- 8. E. Wolanski, Y. Mazda, P. Ridd, "Mangrove hydrodynamics" in Tropical Mangrove Ecosystems, A. I. Robertson, D. M. Alongi, Eds. (Coastal and Estuarine Studies, American Geophysical Union, 2013), vol. 41, pp. 43-62.
- 9. A. McIvor, I. Moller, T. Spencer, M. Spalding, "Reduction of wind and swell waves by mangroves" (Natural Coastal Protection Series Rep. 1, The Nature Conservancy, 2012).
- 10. K. Katija, J. O. Dabiri, A viscosity-enhanced mechanism for biogenic ocean mixing. Nature 460, 624-626 (2009).
- 11. M. E. Huntley, M. Zhou, Influence of animal on turbulence in the sea. Mar. Ecol. Prog. Ser. 273, 65-79 (2004).
- 12. H. U. Riisgård et al., The exhalant jet of mussels Mytilus edulis. Mar. Ecol. Prog. Ser. 437, 147-164 (2011).
- 13. E. Kunze, Biologically generated mixing in the ocean. Annu. Rev. Mari. Sci. 11, 215-226 (2019).
- 14. E. C. Reid et al., Internal waves influence the thermal and nutrient environment on a shallow coral reef. Limnol. Oceanogr. 64, 1949-1965 (2019).
- N. L. Jones, S. G. Monismith, The influence of whitecapping waves on the vertical structure of turbulence in a shallow estuarine embayment. J. Phys. Oceanogr. 38, 1563-1580 (2008)
- 16. B. Kjerfve, J. A. Proehl, Velocity variability in a cross-section of a well-mixed estuary. J. Mar. Res. 37, 409–418 (1979).
- 17. B. S. Holland, M. N. Dawson, G. L. Crow, D. K. Hofmann, Global phylogeography of Cassiopea (Scyphozoa: Rhizostomeae): Molecular evidence for cryptic species and multiple invasions of the Hawaiian Islands. Mar. Biol. 145, 1119-1128 (2004).
- 18. J. Fleck, W. K. Fitt, Degrading mangrove leaves of Rhizophora mangle (Linne) provide a natural cue for settlement and metamorphosis of the upside down jellyfish Cassiopea xamachana (Bigelow). J. Exp. Mar. Biol. Ecol. 234, 83-94 (1999).
- W. Niggl, C. Wild, Spatial distribution of the upside-down jellyfish Cassiopea sp. 19. within fringing coral reef environments of the Northern Red Sea: Implications for its life cycle. Helgol. Mar. Res. 64, 281-287 (2010).
- 20. A. C. Morandini, S. N. Stampar, M. M. Maronna, F. L. Da Silveira, All non-indigenous species were introduced recently? The case study of Cassiopea (Cnidaria: Scyphozoa) in Brazilian waters. J. Mar. Biol. Assoc. U. K. 97, 321-328 (2016).
- 21. S. J. Keable, S. T. Ahyong, First records of the invasive "upside-down jellyfish", Cassiopea (Cnidaria: Scyphozoa: Rhizostomeae: Cassiopeidae), from coastal lakes of New South Wales, Australia. Record Aust. Mus. 68, 23-30 (2016).
- 22. E. Özgür, B. Öztürk, A population of the alien jellyfish, Cassiopea andromeda (Forsskål, 1775) (Cnidaria: Scyphozoa: Rhizostomea) in the Ölüdeniz Lagoon, Turkey. Aquat. Invasions 3, 423-428 (2008).
- 23. E. W. Stoner, C. A. Layman, L. A. Yeager, H. M. Hassett, Effects of anthropogenic disturbance on the abundance and size of epibenthic jellyfish Cassiopea spp. Mar. Pollut. Bull. 62, 1109-1114 (2011).
- 24. C. L. Hamlet, L. A. Miller, T. Rodriguez, A. Santhanakrishnan, "The fluid dynamics of feeding in the upside-down jellyfish" in Natural Locomotion in Fluids and on Surfaces: Swimming, Flving, and Sliding, S. Childress, A. Hosoi, W. W. Schultz, J. Wang, Eds. (IMA Volumes in Mathematics and Its Applications, Springer, 2012), vol. 155, pp. 35–51.
- 25. A. Santhanakrishnan, M. Dollinger, C. L. Hamlet, S. P. Colin, L. A. Miller, Flow structure and transport characteristics of feeding and exchange currents generated by upsidedown Cassiopea jellyfish. J. Exp. Biol. 215, 2369-2381 (2012).
- 26. C. Jantzen, C. Wild, M. Rasheed, M. El-Zibdah, C. Richter, Enhanced pore-water nutrient fluxes by the upside-down jellyfish Cassiopea sp. in a Red Sea coral reef. Mar. Ecol. Prog. Ser. 411, 117-125 (2010).
- 27. J. M. Morell, J. E. Corredor, Sediment nitrogen trapping in a mangrove lagoon. Estuarine Coast. Shelf Sci. 37, 203-212 (1993).
- 28. A. M. Lohrer, N. J. Halliday, S. F. Thrush, J. E. Hewitt, I. F. Rodil, Ecosystem functioning in a disturbance-recovery context: Contribution of macrofauna to primary production and nutrient release on intertidal sandflats. J. Exp. Mar. Biol. Ecol. 390, 6-13 (2010).
- 29. D. M. Durieux, K. T. Du Clos, B. J. Gemmell, Morphological measurements. FigShare. https://figshare.com/articles/dataset/Morphological_Measurements_KML_2017/9700406. Deposited 21 August 2019.
- 30. D. M. Durieux, K. T. Du Clos, B. J. Gemmell, KML 2016 population transects. FigShare. https://figshare.com/articles/dataset/KML_2016_Population_Transects/9700370. Deposited 21 August 2019.

- 31. D. M. Durieux, K. T. Du Clos, B. J. Gemmell, PIV videos of Cassiopea vertical jets. Figshare. https://figshare.com/articles/dataset/PIV_videos_of_Cassiopea_vertical_jets/ 9753149. Deposited 30 August 2019.
- 32. D. M. Durieux, K. T. Du Clos, B. J. Gemmell, PIV videos of clusters of Cassiopea. FigShare. https://figshare.com/articles/dataset/PIV_videos_of_clusters_of_Cassiopea_ vertical_jet_/9754916. Deposited 30 August 2019.
- 33. D. M. Durieux, K. T. Du Clos, B. J. Gemmell, Interstitial Porewater Release. FigShare. https://figshare.com/articles/dataset/Interstitial_Porewater_Release/9701534. Deposited 21 August 2019.
- 34. L. A. Leonard, M. E. Luther, Flow hydrodynamics in tidal marsh canopies. Limnol. Oceanogr. 40, 1474-1484 (1995).
- 35. B. K. Norris, J. C. Mullarney, K. R. Bryan, S. M. Henderson, The effect of pneumatophore density on turbulence: A field study in a sonneratia-dominated mangrove forest, Vietnam. Continent. Shelf Res. 147, 114-127 (2017).
- 36. B. T. Kannan, N. R. Panchapakesan, Effect of momentum flux distribution on multiple round jets. Aircraft Eng. Aero. Technol. 90, 452–460 (2018).
- 37. K. R. Sutherland, D. Weihs, Hydrodynamic advantages of swimming by salp chains. J. R. Soc. Interface 14, 20170298 (2017).
- 38. R. Mann, M. Southworth, J. M. Harding, J. A. Wesson, Population studies of the native eastern oyster, Crassostrea virginica, (Gmelin, 1791) in the James River, Virginia, USA. J. Shellfish Res. 28, 193-220 (2009).
- 39. R. I. E. Newell, E. W. Koch, Modeling seagrass density and distribution in response to changes in turbidity stemming from bivalve filtration and seagrass sediment stabilization. Estuaries 27, 793-806 (2004).
- 40. C. Mo, B. Neilson, Standardization of ovster soft tissue dry weight measurements. Water Res. 28, 243-246 (1994).
- 41. D. J. H. Phillips, The common mussel Mytilus edulis as an indicator of pollution by zinc, cadmium, lead and copper. I. Effects of environmental variables on uptake of metals. Mar. Biol. 38, 59-69 (1976).
- 42. P. J. Dare, D. B. Edwards, Seasonal changes in flesh weight and biochemical composition of mussels (Mytilus edulis L.) in the Conwy Estuary, North Wales. J. Exp. Mar. Biol. Ecol. 18, 89-97 (1975).
- 43. K. T. Du. Clos, I. T. Jones, T. J. Carrier, D. C. Brady, P. A. Jumars, Model-assisted measurements of suspension-feeding flow velocities. J. Exp. Biol. 220, 2096-2107 (2017)
- 44. J. K. Petersen, I. Svane, Filtration rate in seven Scandinavian ascidians: Implications of the morphology of the gill sac. Mar. Biol. 140, 397–402 (2002).
- 45. S. McGrorty, R. T. Clarke, C. J. Reading, J. D. Goss-Custard, Population dynamics of the mussel Mytilus edulis: Density changes and regulation of the population in the Exe estuary, Devon. Mar. Ecol. Prog. Ser. 67, 157-169 (1990).
- 46. I. Svane, Ascidian reproductive patterns related to long-term population dynamics. Sarsia 68, 249-255 (1983).
- 47. E. Carrington, G. M. Moeser, S. B. Thompson, L. C. Coutts, C. A. Craig, Mussel attachment on rocky shores: The effect of flow on byssus production. Integr. Comp. Biol. 48, 801-807 (2008).
- 48. L. A. Van Duren, P. M. J. Herman, A. J. J. Sandee, C. H. R. Heip, Effects of mussel filtering activity on boundary layer structure, J. Sea Res. 55, 3–14 (2006).
- 49. E. Wolanski, Sediment transport in mangrove swamps. Hydrobiologia 295, 31-42 (1995)
- 50. T. M. Gihring, A. Canion, A. Riggs, M. Huettel, J. E. Kostka, Denitrification in shallow, sublittoral Gulf of Mexico permeable sediments. Limnol. Oceanogr. 55, 43-54 (2010).
- 51. R. Rosenberg, Marine benthic faunal sucessional stages and related sedimentary activity. Sci. Mar. 65 (Suppl. 2), 107-119 (2001).
- 52. A. M. Szmant, A. Forrester, Water column and sediment nitrogen and phosphorus distribution patterns in the Florida Keys, USA. Coral Reefs 15, 21-41 (1996).
- 53. C. T. Rueden et al., Imagej2: Imagej for the next generation of scientific image data. BMC Bioinf. 18, 529 (2017).
- 54. J. Schindelin et al., Fiji: An open-source platform for biological-image analysis. Nat. Methods 9, 676-682 (2012).
- 55. S. B. Pope, "Free shear flows" in Turbulent Flows, S. B. Pope, Ed. (Cambridge University Press, 2000), pp. 96-134.
- 56. J. De Jong et al., Dissipation rate estimation from piv in zero-mean isotropic turbulence. Exp. Fluid. 46, 499-515 (2009).
- 57. D. M. Durieux, K. T. Du Clos, B. J. Gemmell, Lido Key ADV Data. FigShare. https://figshare.com/articles/dataset/Lido_Key_ADV_Data/9696272. Deposited 20 August 2019.

